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# UNITED STATES DEPARTMENT OF THE INTERIOR

# TOPOGRAPHIC CHARACTERISTICS OF DRAINAGE BASINS

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# UNITED STATES DEPARTMENT OF THE INTERIOR

J. A. Krug, Secretary

GEOLOGICAL SURVEY W. E. Wrather, Director

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# TOPOGRAPHIC CHARACTERISTICS OF DRAINAGE BASINS

By WALTER B. LANGBEIN
AND OTHERS

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#### TOPOGRAPHIC CHARACTERISTICS OF DRAINAGE BASINS

# By WALTER B. LANGBEIN and others

#### ABSTRACT

River floods are the result of many causes, and one of the primary objectives of scientific hydrology is the segregation and evaluation of the causative factors. The climatic factor and the soil-vegetation complex are variables that exercise their principal influence on the volume of runoff. The topography of drainage basins is a sensibly permanent characteristic which influences mainly the concentration or time distribution of the discharge from a drainage basin. River systems differ in their efficiency as agencies for collecting and conducting water. In some systems, surface waters are quickly assembled, and the discharge reflects somewhat sensitively the variations of the available supply. In others, the surface drainage is longer delayed and the discharge is released slowly.

As a basis for quantitative studies of these evident differences in behavior, selected topographic features for about 340 drainage basins in the northeastern United States were studied, using Geological Survey topographic maps. The data were compiled in cooperation with the Work Projects Administration of the Federal Works Agency and included information on drainage area, length of streams, stream density, land slope, channel slope, area-altitude distribution, and area of water bodies of basins that ranged in extent from 1.64 to 7,797 square miles. Considerable effort was made to assure accuracy of the computations by appropriate checks, and the results are summarized in the table at the end of this report.

The results indicate that none of the topographic factors are unique, but each reflects a condition that also influences the others. For example, steep land slopes are generally associated with steep channel slopes and conversely. A significant variation of slope and altitude with area of basin is found, and stream density tends to vary with the land slope.

### INTRODUCTION

This report presents a compilation of topographic data on drainage basins in the northeastern United States. The configuration of the earth reflects the impact of many natural forces, and it in turn exercises profound influence upon man. Most of these influences are so basic that they have shaped life and civilization into conformity with them. Mountains, plains, valleys, and rivers each favor or retard man's search for economic stability. Within human history the first three have remained unchanged. Rivers, on the other hand, fluctuate in size from day to day and from year to year. The amplitude and frequency of these fluctuations, so significant with respect to navigation, water power, irrigation, and such riparian developments as cities and highways, are largely determined by three separate, yet interdependent features, namely climate, physiography, and the soil-vegetation complex. The interrelation of these three features with the behavior of rivers is imperfectly understood and is the subject of much investigation. This report singles out the physiography of the land for attention.

The relations between the rate, volume, and fluctuations of rivers and the topographic characteristics of the land they drain and through which they flow may be readily determined after discerning examination of the terrain and river developments, but expressing them in the quantitative terms necessary for the economic design of structures for river utilization or control requires first, topographic maps, and second, records of river flow of length adequate to define the behavior.

The stream-gaging program of the Geological Survey is Nation-wide and now includes over 4,500 river-measurement stations, at which more than 65,000 station years of record were available in 1942. These records furnish an adequate source of material concerning stream behavior. The mapping program of the Geological Survey, also Nation-wide, is in general not so complete. Although about 50 percent of the country has been mapped, only States in the north-eastern part have been completely covered; the scattered areas mapped in other States generally do not cover completely the areas in which stream-gaging has been carried on, so that only a small fraction of them are suitable for use in comparisons of stream-flow characteristics or river morphology.

In the northeastern and north-central States the range in topography is sufficient to furnish a basis for studying its effect on stream flow. The topographic characteristics compiled from the maps and summarized in this report can only be evaluated by a consideration of the hydrology of stream flow, the assembling of waters in a drainage system, and the hydraulic elements that regulate velocity of flow. Many stream-flow characteristics are related either directly or indirectly to topographic features. It would seem, however, that the factors most sensitive to topographic difference would be those relating to floods. In this study, therefore, particular although not exhaustive attention is given to the correlation of flood-flow characteristics with topography. This information will serve as a basis for

further study of such correlations and also of other characteristics, such as volume yield, erosion, and deposition of sediments. Similarly the topographic data offers basic material for studies of river morphology, as geologic evidence suggests that a significant portion of river-channel development takes place during flood.

# COOPERATION AND PERSONNEL

The cooperative project for the compilation of topographic data was undertaken in 1939 by the Works Progress Administration, which on April 25, 1939, became the Work Projects Administration under the Federal Works Agency. Their cooperation in organizing competent working groups is especially acknowledged. The Geological Survey sponsored the project and furnished technical direction, maps, and supplies. This work was carried on by W. B. Langbein, under the general direction of R. W. Davenport, chief of the Division of Water Utilization. The project at Boston, Mass., was under the supervision of H. B. Kinnison, district engineer, and his associates, particularly C. E. Knox, M. A. Benson, and B. R. Colby. duct of the work at Pittsburgh, Pa., was ably managed by Wm. S. Crozier, supervisor for the Works Projects Administration. Crozier died January 21, 1941, near the close of the project. Erskine, associate engineer of the Geological Survey at Pittsburgh, maintained close contact with the project there, and its continuity and efficiency may be largely credited to his competent administration.

## METEOROLOGIC FACTORS AFFECTING RUNOFF

River floods are the result of many causes, and one of the primary objectives of hydrologic study is the segregation of the causative factors and the evaluation of their effects on the resultant floods under various associated conditions.

Readily apparent is the source of the water, generally an unusual amount of rainfall, which may be characterized by great intensity and in many regions may be augmented by water from melting snow. Water in excess of that which can be absorbed by the ground or evaporated into the air directly or through vegetation collects in the stream channels that drain the area. Once in the stream system, the runoff flows to the mouth through channels which, as the trunk of a network of streams, progressively increase in size as contributions are received from tributary streams.

The quantity of rainfall or snow melt, its time distribution, and the associated soil, vegetal, and climatic conditions that determine the portion of the supply that becomes direct runoff are to a large extent variable characteristics of individual storms. These variable edaphic and climatic factors are separate phases of the rainfall-runoff relation.

The channel system, however, is a relatively permanent characteristic of a drainage basin. Some influential changes may take place in this system; for example, variations in seasonal vegetation along the banks may affect the hydraulic conveyance, floods may scour or deposit sediments, and old bends may be cut through and new ones created. Although the effect of these changes on local flood stages may be considerable, it is assumed that the resultant effect on total discharge from a basin will tend to be compensating and that even the cumulative effect on flood-discharge characteristics during a period as short as the usual stream-gaging record would be minor.

# DIFFERENCES IN CHARACTER OF DRAINAGE BASINS

River systems differ in their efficiency as agencies for collecting and conducting water. In some systems the surface waters are quickly assembled, and the discharge therefrom reflects somewhat sensitively the variations of the available supply. In others the surface drainage is longer delayed and the discharge is released slowly. This difference is illustrated in figure 48, which shows the hydrographs of two nearby streams, each draining about 50 square miles of coastal areas of New Jersey, during a flood in June 1938. The rainfall causing these floods and the volume discharged were nearly the same for both areas. The difference in behavior illustrated by the hydrographs is normal for these two basins and may be accounted for largely by the differences in physiographic characteristics, Manasquan River having about twice the gradient of Great Egg River and about one-fifth the swamp area.

# PREVIOUS STUDIES

The fact of relationship between the time distribution of discharge during a flood and the size, shape, and gradient of a drainage system is widely recognized. Few attempts have heretofore been made to determine this relationship quantitatively, probably because of the volume of labor required to evaluate the topographic factors.

Horton in 1926 and again in 1932 discussed the desirability and need for a quantitative rational procedure and proposed methods for evaluating certain pertinent physiographic factors.

Pettis <sup>2</sup> in 1927 presented a formula to compute the maximum flood discharge, in which the five-fourths power of the average basin width was used.

Gregory and Arnold 3 in 1932 developed in detail certain expressions

<sup>&</sup>lt;sup>1</sup> Horton, R. E., in Jarvis, C. S., Flood'flow characteristics: Am. Soc. Civil Eng. Trans., vol. 89, pp. 1681-1086, 1926; Drainage-basin characteristics: Am. Geophys. Union Trans., No. 13, pp. 350-361, 1932.

<sup>&</sup>lt;sup>2</sup> Pettis, C. R., A new theory of river flood flow (published privately, copyrighted 1927).

<sup>&</sup>lt;sup>2</sup> Gregory, R. L., and Arnold, C. E., Rational runoff formulas: Am. Soc. Civil Eng. Trans., vol. 96, pp. 1038-1175, 1932.

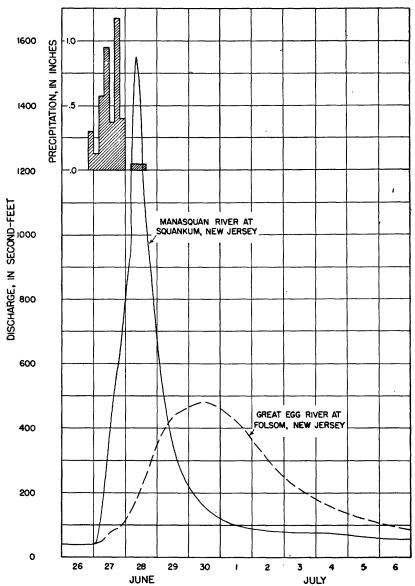


FIGURE 48.—Hydrographs of two streams in New Jersey during flood of June 1938.

and procedures applicable to small drainage areas for translating rainfall into rates of stream flow in terms of basin characteristics.

Bernard <sup>4</sup> carried Gregory and Arnold's expressions somewhat further and presented formulas applicable to a few selected basins.

<sup>&</sup>lt;sup>4</sup> Bernard, M. M., An approach to determinate stream flow: Am. Soc. Civil Eng. Trans., vol. 100, pp. 347-395, 1935.

<sup>747049-47-2</sup> 

Sherman <sup>5</sup> in 1932 presented unit hydrographs for four streams whose drainage areas and slopes differed widely. He explained how the unit hydrographs expressed these differences and suggested that basins having physical characteristics similar to the four types presented would have similar hydrographs.

McCarthy <sup>6</sup> in 1937 also used the unit hydrograph as an expression of the runoff characteristics of a drainage basin, stating further that "the agreements between graphs developed from May and November storms substantiate the contention that primarily the unit hydrograph is a function not of surface cover, which may be subject to seasonal change, but of topographic features of a watershed." From this postulate he derived a working relationship between the crest discharge and the length of base of unit hydrographs of 25 drainage basins in the Connecticut River Basin, in terms of the area, mean slope, and stream pattern, determined by inspection and expressed as one-stem basin, two-stem basin, and so on, for application to flood-control design.

Morgan and Hullinghorst 7 in 1939 stated:

The factors which determine the discharge characteristics of any watersned are innumerable, some having a major bearing on those characteristics while others are of negligible consequence. It was determined by examination of nine gaged basins having complete unit hydrograph and watershed data, and corroborated by examination of a number of gaged basins with data ranging from almost complete to fragmentary, that the discharge characteristics can be attributed principally to three fundamental, definite watershed characteristics, namely,

- a. Area of the watershed in square miles,
- b. Mean length of travel in miles, and
- c. Mean height of watershed above outflow station in feet.

On this basis, empirical relations between these three factors were established for nine streams tributary to the Chemung River in New York.

#### PURPOSE AND SCOPE OF THE PRESENT STUDY

The present project was designed to provide basic material whereby investigations such as those outlined can be carried further, the range being limited, of course, to areas adequately mapped, which are mainly in the northeastern United States. Prior to this study, references to the subject were read with the view to determining which topographic factors were considered to have major influence upon

Sherman, L. K., The relation of hydrographs of runoff to size and character of drainage basins: Am. Geophys. Union Trans., No. 13, pp. 332-339, 1932.

<sup>6</sup> McCarthy, G. T., The unit hydrograph and flood routing (unpublished manuscript presented at conference of North Atlantic Division, Corps of Engineers, U. S. Army, June 24, 1938).

<sup>&</sup>lt;sup>7</sup> Morgan, R., and Hullinghorst, D. W., Unit hydrographs for gaged and ungaged watersheds: U. S. Engineer Office, Binghamton, N. Y., July 1939. [Processed.]

discharge characteristics, and, so far as practicable, the suggestions thus obtained were incorporated in the project as proposed for cooperation to the Works Progress Administration (succeeded on April 25, 1939, by the Work Projects Administration). The compilation was based upon the topographic maps of the Geological Survey covering areas tributary to gaging stations of the Geological Survey.

In the organization of a surface-water system, and of a large part of the ground-water system as well, the drainage basin is a natural hydrologic land unit. Surface runoff is divided into drainage basins by the watersheds, and within each basin it follows a system of water courses in which the flow undergoes retardation, acceleration, or other changes that are distinctly related to the physical characteristics of that basin. Similar conditions exist with respect to all or most of the ground-water runoff. Essentially all the water within a given basin, except that which is lost by evaporation or transpiration, drains out through a common outlet or mouth.

For purposes of analysis, a major stream basin may be subdivided by considering the area tributary to the stream at any given point, for example, a gaging station, as a basin having its own characteristics. The separate characteristics of several contributory areas may then be combined to obtain the resultant for the major basin.

Geographic and topographic characteristics of drainage basins, based largely on certain horizontal and vertical dimensions, were selected for compilation and study. Geographic characteristics include water bodies, direction of stream flow, latitude, and longitude. Topographic characteristics include horizontal dimensions covering basin area, stream length, and area-distance distribution, and vertical dimensions covering land slope, tributary and principal stream slopes, and basin altitude.

In selecting basins for this study preference was given to those for which long-term stream-flow records are available and to those free from artificial regulation. In addition to areas in the northeastern States and the Ohio River Basin a few surveyed areas in Wisconsin and Kentucky were included to spread the range in geographic extent and topographic characteristics. Many basins in New York, Pennsylvania, and New Jersey that might otherwise have been included were not studied because of insufficient time.

Besides listing topographic and hydrologic data, the original records afford a gazetteer of streams and lakes. Maps were prepared showing the stream skeleton of each basin, with names of streams, length from confluence to confluence, and drainage areas and altitudes pertinent to the subdivisions. A list of lakes and ponds, giving names, locations, and approximate altitudes and areas was also prepared, much of the

data for basins in New York State being based on a gazetteer by E. M. Douglas.<sup>8</sup>

The summarized results of the compilation, covering about 340 basins, are given in the table on pages 145-155. The original records are on file at the office of the Geological Survey in Washington, D. C., and the computations for basins in New England (except Maine) are on file also at the Boston office of the Geological Survey.

#### METHODS OF WORK

#### MAPS

Quadrangle maps on the scale 1:62,500, are the basis of this compilation, except for a few areas in New Jersey where more detailed maps were available. On these maps the gaging stations were located and the tributary basins with sub-basin divisions were outlined. Generally, each basin was divided into 50 to 75 sub-basins. Care was taken that the sub-basins crossed the streams only at confluence points. To systematize the necessary tabulations, the sub-basins were numbered in accord with the following system, which is illustrated on plate 2.

The headwater basin farthest upstream (the one farthest removed from the gaging station along the main stream) is called no. 1; the sub-basin which it joins at the first confluence point is called no. 2; the sub-basin (or intervening area along the combined channel) below the confluence of sub-basins 1 and 2 is called no. 3; the next tributary sub-basin is called no. 4, and so on. Where a large tributary stream that has been subdivided joins the main stream, the next consecutive number is assigned to the farthest upstream sub-basin of this tributary. This constitutes the lowest number on such tributary, and the sub-basins of the tributary system are then numbered in the same manner as those of the main stream, down to the sub-basin immediately above the confluence of the tributary with the main stream. The succeeding number is assigned to the sub-basin along the main river immediately below the confluence, as before. The highest number in a basin is that of the sub-basin immediately above the mouth of the main stream: it indicates the number of sub-basins into which that drainage area is divided. As shown on plate 2 the number of each sub-basin is the large integer near its center; the smaller figures represent area, length of stream, and altitude.

<sup>&</sup>lt;sup>8</sup> Douglas, E. M., Gazetteer of the lakes, ponds, and reservoirs of the State of New York: Map Information Office, Board of Surveys and Maps, 44 pp., Washington, 1926. [Processed.]

#### AREA OF BASINS

The total area of the basin within the watershed lines above the selected gaging station is the primary basin factor. In a humid climate the volume of discharge varies directly with the size of the tributary drainage area. Accordingly, the area in square miles was measured, not only of the main basin above the gaging station but also of a number of sub-basins (generally over 50). The size of the basins included in this compilation ranges from 1.64 to 7,797 square miles. In general, large streams were excluded because the size of sheets became unwieldy, because some contained unsurveyed areas, and because their stream-flow characteristics could best be determined by synthesis of their components.

## STREAM DENSITY

The runoff from the several parts of the drainage basin is discharged by the streams, and, other factors being constant, the time required for the water to flow a given distance is directly proportional to the stream length. The stream or drainage density is the ratio between the total length of all streams within the drainage basin and the total area of the basin and is an indication of the drainage development. Accordingly, the length of all streams down to the smallest shown on the topographic maps was measured to determine the stream density and the area-distance distribution.

The number of small headwater streams shown on the topographic maps would vary with the season and the wetness of the particular year during which the survey was made, as well as with the judgment of the topographer and cartographer as to the amount of detail to be shown on the map. These circumstances introduce a measure of inconsistency in stream-density results as determined from maps.

The ratio of stream density for the basins included in this compilation, all of which are in the humid region, ranges from 0.89 to 3.37 miles per square mile and averages 1.65 miles per square mile. Other factors being equal, high drainage density indicates a more effective operation, of the agencies of stream incision. Greater incision, for example, would be associated with steep land slopes. Opportunity for incision would be greater also where most of the discharge occurs as surface runoff rather than through ground-water channels; such a condition exists in areas where the ground is sufficiently impervious to shed storm rainfall. Drainage density is greater in humid regions than in arid regions; it would approach zero in flat, sandy desert regions and would approach a maximum in steep, rocky, humid regions.

The variation of stream density with the land slope is shown by the following data derived from groups of basins in New England.

# Variation of stream density with land slope

Range in stream density (miles per square mile)	A verage land slope (feet per mile)	
	290	,
1.26 to 1.50	550	
,	600	
2.01 to 2.25	700	

The mean land slopes for basins in New England having drainage densities within the ranges indicated in the above data were averaged, and the results indicate that, in general, in a given region the higher drainage densities are associated with the steeper land slopes. The reciprocal of the drainage density is the average distance between streams, and half the reciprocal of drainage density is the average horizontal distance between the streams and appurtenant watershed lines, measured at right angles to the streams. Drainage density appears to be inversely related to the distance of overland flow as distinguished from flow in stream channels. However, in basins sufficiently permeable so that all rainfall can be taken directly into the soil through infiltration, the drainage density approaches zero and is associated with zero overland flow.

#### AREA-DISTANCE DISTRIBUTION

The concentration of runoff from drainage basins of equal size may be greatly affected by the distribution of the area with respect to distance from the gaging station or outlet. Other factors being equal, the runoff from areas close to the gaging station should reach it sooner than water from remote areas. Accordingly, a drainage basin whose tributaries are compactly organized, so that water from all parts of the basin has a comparatively short distance to travel, will discharge its runoff more quickly and reach greater flood crests than one in which the larger part of the area is remote from the gaging station or outlet. This basin characteristic is expressed in the summary table by the quantity  $\Sigma$ al, computed by multiplying each partial area in the basin (a) by the channel distance from the midpoint of the main stem serving it, downstream to the gaging station (l). Distances along the stream channels were measured in 0.1 mile chords.

In a sense this quantity is also a measure of the volume of channel storage in the basin. For example, if under a given regime of flow the cross-sectional area of a river at a given place varies directly as the drainage area above, then the volume in any given reach would vary as the product of a coefficient by mean drainage area above the reach by the length of the reach. The coefficient would be a function of the stage of the regime selected, the slope of the reach, frictional resistance, and other hydraulic factors. No method is proposed for evaluating the coefficient. However, the sum of the products of mean drainage area and length of reaches for a given basin is equal to the product  $\Sigma$ al which was derived by both methods of computation.

The most compact drainage basin would be a glory-hole inlet, and the product Σal for such a basin is 0.375 A <sup>1.50</sup> where A represents total area; for an equilateral triangle, with reference to an outlet at one of the vertices, the product is 0.94 A <sup>1.50</sup>; and for a square, with reference to a corner, it is 0.76 A <sup>1.50</sup>. Figure 49 shows the results of plotting the products Σal against the corresponding drainage area. Only enough points are shown to define the line of regression, whose equation is 0.90 A <sup>1.56</sup>, or more approximately 1.2 A <sup>1.50</sup>, within the range shown. Natural basins are generally less compact than any of the geometric shapes mentioned.

Additional subdivision of a basin beyond the 50 to 75 sub-basins generally used would tend to increase the value of the product Σal. However, a study of West River Basin, above Newfane, Vt., indicates that the product Σal for 20 sub-basins was 6,620, for 50 sub-basins 6,810, and for 100 sub-basins 6,860. The values given in this report may therefore be considered essentially correct limiting values.

Points on the right of the trend line (fig. 49) represent basins less compact than the average, and those on the left the more compact. The regression line therefore furnishes a standard for comparing the relative compactness of different basins.

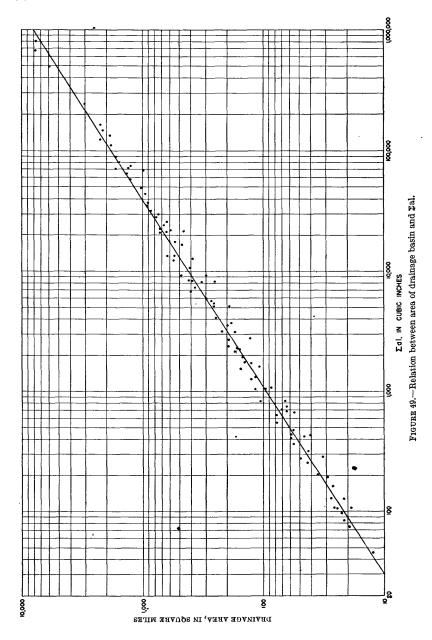
#### LENGTH OF BASIN

The table (pp. 145-155) lists the length of longest watercourse in each basin measured in 0.1-mile chords to the source of the most headward stream. This length, when divided by the mean velocity of flow will give the time of concentration as used in the rational formula for the computation of flood discharge.

The mean length of travel of runoff or the distance to the center of gravity of the drainage system may be found by dividing the quantity  $\Sigma$ al by the drainage area in square miles. This quotient is commonly identified by the symbol  $L_{ca}$ . The table also lists the length of principal streams as defined under "Channel slope."

# LAND SLOPE

Rainfall or snow melt which becomes direct runoff flows over the surface of the ground or, where the surface soil is shallow and permeable, immediately beneath it over the bedrock. The average distance water travels before entering a stream channel may be



expressed in terms of stream density. (See p. 133.) The rapidity with which the water travels to the streams likewise depends on the slope of the land.

The contours on the topographic maps provide a basis for determing the slope of the land by the intersection-line method outlined by Horton <sup>9</sup> as follows:

The intersection-line method.—In order to reduce the labor of computation of slope of large areas the author has utilized the following method. An area the slope of which is to be determined is subdivided into squares of equal size by lines forming the boundaries between adjacent squares. The number of contours crossed by each subdividing line is counted and the lengths of the lines are scaled. Then the average scale-distance l' between contour crossing in the subdivision lines is

$$l' = \frac{\Sigma l}{N}$$

where N is the number of contours crossed and  $\Sigma l$  is the total length of the subdividing lines. If  $\alpha$  is the horizontal angle at which each of two parallel contours crosses an intersection line, then l' sin  $\alpha$  is the horizontal distance between the two contours measured normal to the contours. Contours may cross the intersection lines at all angles from  $0^{\circ}$  to  $90^{\circ}$ . The mean value of  $\sin \alpha$  for angles from  $0^{\circ}$  to  $90^{\circ}$  is

$$\int_0^{\frac{\pi}{2}} \frac{\sin \alpha d\alpha}{\pi} = \frac{2}{\pi} = 0.6366$$

If D is the contour interval or difference in elevation in feet, and L is the average normal horizontal distance between contours, then

$$L=0.6366 l'$$

and the mean slope Sg of the area is

$$Sg = \frac{D}{0.6366 \frac{\Sigma l}{N}}$$

=1.571 
$$\frac{DN}{\Sigma l}$$

In applying this method it is assumed that each contour crossed represents a difference of elevation along the subdivisional line equal to the contour interval. Of course it may happen that two adjacent contours are at the same elevation and are separated by land only a little higher or lower. On an average, however, the elevations of summits or depressions between equal contours will differ from that of the adjacent contours by an amount equal to one-half the contour interval, and it can readily be seen that the average declivity between a pair of contours of equal elevation is nearly the same as if the contours were separated by the contour interval D, so that the method gives nearly correct results even where the subdivision lines cross adjoining contours of equal elevation, as in the case of summits and depressions.

Horton, R. E., Drainage-basin characteristics: Am. Geophys. Union Trans., No. 13, pp. 350-361, 1932. 747049-47---3

By making the subdivision lines sufficiently frequent, the average slope of an area may be determined with whatever degree of accuracy is required.

This method has been tested by comparison of slope for the same area computed from the measured total lengths of contours, with, in general, good agreement.

In carrying out this computation, the slope along the meridian lines is computed separately from the slope along the parallels of latitude. Where there is a great difference between the land slope in the two directions, the orientation of the basin is determined by the axis of least slope. Where the east-west slope and the north-south slope are nearly the same, the line of orientation may be approximately midway between the two, or it may not be clearly defined in either direction in a cup-shaped or fan-shaped basin. Land slopes listed in the table range from 1,598 feet per mile for the upper Pemigewasset River Basin in New Hampshire to 55 feet for Great Egg River in Coastal New Jersey.

Paulsen <sup>10</sup> found, during the flood of September 1938 in the North Atlantic States, that the infiltration index tends to increase with decrease in mean land slope. He states that "although the slope of the land itself might influence the retentive capacity of the ground, this tendency may be due to other factors related to slope, such as depth of soil cover."

#### CHANNEL SLOPE

Upon leaving the land the runoff enters the channel system, through which it flows in channels that increase progressively in size with the entrance of additional water. Channels in a drainage basin are classified for study as principal and tributary. The principal streams of a basin are defined as those that drain 10 percent or more of the total area of the basin; tributaries are defined as those that drain less than 10 percent of the area of the basin. The average slope of the tributaries and of the principal streams is computed separately as the quotient of the total fall divided by the corresponding total length and is reported in the summary table.

In computing the slope of the stream channels, only the largest stream in each sub-basin is considered. Thus, if a basin is divided into 75 sub-basins only 75 stream lengths and falls are measured. These stream lengths are classified as principal or tributary, and the average slope of each is computed. As only one stream in each sub-basin is included in the classification, many minor headwater streams are excluded from consideration; consequently, the reported slope of the tributary stream is affected by the number of subareas into which the basin is divided. The reported mean slope of the tributary

<sup>&</sup>lt;sup>10</sup> Paulsen, C. G., Hurricane floods of September 1938: U. S. Geol. Survey Water-Supply Paper 867, pp. 440-441, 1940.

streams increases as the number of subareas becomes larger, thus embracing more steep minor headwater streams. This is illustrated in figure 50, which shows the result of a comprehensive study of the slope of tributaries of West River at Newfane, Vt. The asymptote resulting from that study is about 225 feet per mile, whereas the channel slope obtained with 53 subareas (see table, No. 1–354) is 200 feet per mile.

The slope information for the several drainage basins listed in the table discloses that a steep land slope is generally associated with

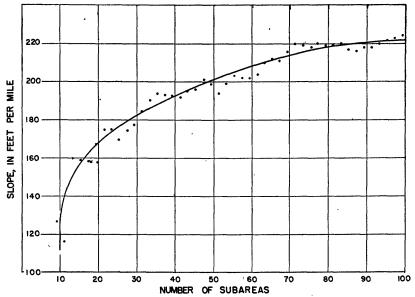


FIGURE 50.—Variation of computed slope of tributary streams with number of subareas, West River at Newfane. Vt.

steep tributary and principal channel slopes and conversely, as might be expected. There is, however, no systematic variation; moreover, according to geometric analysis by Horton,<sup>11</sup> the ratio between principal channel slope and average land slope is a measure of the horizontal angle that the lateral slope makes relative to the stream slope. A low slope ratio indicates that the laterals tend to enter the streams at right angles, whereas the angle of inflow into the stream becomes more acute as the channel slope approaches equality with the ground slope. This slope ratio tends to decrease with increase in drainage area, but varies considerably between drainage basins of equal size.

<sup>11</sup> Horton, R. E., op. cit., p. 360.

#### AREA-ALTITUDE DISTRIBUTION

Another method of expressing the slope of the basin is by means of the altitude of the several parts with reference to sea level. This is best expressed through the hypsometric curve, as a graph showing the area-altitude distribution is called. Although the area-altitude distribution was derived for each basin, only the maximum, mean, and minimum altitudes as determined from the topographic maps are shown in the table. From this information, however, the area-altitude distribution curve can be readily approximated. Figure 51 shows the hypsometric curves for several basins plotted in terms of percent of range in altitude and percent of area above the indicated altitude. The variations are wide, but in general the mean altitude

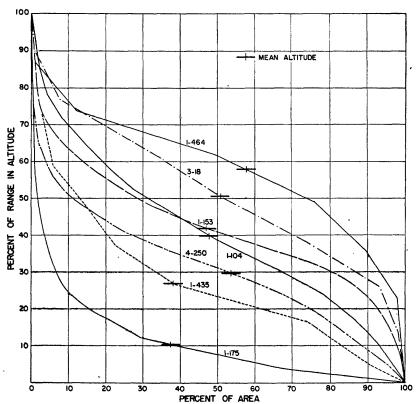


FIGURE 51.-Typical hypsometric curves for drainage basins.

- 1-104. Swift River near Roxbury, Maine.
- 1-153. East Branch of Pemigewasset River near Lincoln, N. H.
- 1-175. Lake Winnipesaukee outlet at Lakeport, N. H.
- 1-435. Quinnipiac River at Wallingford, Conn.
- 1-464. Leadmine Brook near Thomaston, Conn.
- 3- 18. Brokenstraw Creek at Youngsville, Pa.
- 4-250. Otter Creek at Center Rutland, Vt.

of a basin is located at 0.34 of the range between the minimum and maximum; thus a basin is comparable to the surface of a cone.

The area-altitude distribution curve has several applications. For example, snow surveys generally show an increase in depth of cover and water equivalent with increase in altitude; the area-altitude relation provides a means for estimating the mean depth of snow or its water equivalent over a drainage basin. Barrows <sup>12</sup> describes a significant variation in annual precipitation and runoff in the Connecticut River Basin with respect to altitude. The obvious variation in temperature with change in altitude is further indication of the utility of the area-altitude distribution curve.

The mean altitude of the basin above the altitude at the outlet or gaging station represents the potential head of a uniform depth of water over the basin with respect to the outlet or gaging station, and as such is a factor in determining the rate at which the waters are collected and discharged. The data in the summary table shows that, in general, the land slopes and channel slopes vary with the mean altitude of the basin above the outlet. Thus steep slopes are associated with a high altitude above the outlet, and conversely. A rough average relation between slope and mean altitude is as follows:

$$h = K_1 S_1 + K_2 S_t + K_3 S_p$$

where  $K_1 = 0.31$ 

K<sub>2</sub> ranges from 0.97 at 50 square miles to 3.0 at 1,000 square miles.

 $K_3$  ranges from 3.5 at 50 square miles to 23.4 at 1,000 square miles.

 $S_1$ =mean land slope, in feet per mile.

 $S_t$ =slope of tributary streams, in feet per mile.

 $S_p$  = slope of principal streams, in feet per mile.

#### AREA OF WATER SURFACES

The effect of storage in retaining flood runoff and prolonging its discharge until the flood in channels farther downstream has begun to subside tends to reduce flood peaks and increase the time lag between rainfall and its consequent runoff. Natural storage in lakes and ponds and artificial storage in reservoirs aids this retardation. A measure of the amount of storage available for such modification of flood discharge can be derived from the surface area of the water bodies shown on the topographic map. (See table.) The computations at Boston included the determination of swamp areas, which had been part of an earlier project carried on in 1936 in cooperation with the Works Progress Administration; this covered compilations for the Merrimac and Connecticut River Basins, both in square miles and

<sup>&</sup>lt;sup>12</sup> Barrows, H. K., Precipitation and runoff and altitude relations for Connecticut River: Am. Geophys Union Trans., 14th Ann. Meeting, pp. 396-406, 1933.

in drainage percent.<sup>13</sup> The computations made at Pittsburgh did not include swamp areas.

The areas of swamps as reported would be affected by the hydrologic conditions under which the topographic surveys were made. Surveys made in spring or early summer would probably show a greater swamp area than those made in late summer or fall, and surveys in a wet year would show marked contrast with those made in a dry year. It is not known to what extent the results given in the table were affected by hydrologic conditions.

It should be pointed out that the area of water surfaces is only one measure of their effect on the time distribution of flood discharge. The position of the water bodies in the river system is also important; thus a large pond near the headwaters would affect but a small part of the runoff, whereas one of the same size farther downstream would affect a larger part of the runoff.

In addition to the effect of storage in modifying the shape of flood waves or the time distribution of runoff, the total volume of runoff may be influenced by evaporation from lakes, reservoirs, and swamps. The loss of water by evaporation from water surfaces in the northeast is about twice that from land surfaces, per unit of area. Accordingly, basins with a large proportion of water and swamp surfaces may be expected to yield less runoff than those with a small proportion.

In many of the basins listed in the table the proportion of lake and swamp areas exceeds 10 percent, and in a few, especially in New England, it approaches 20 percent; doubtless the effect on water losses is significant. The percentage of lake area is highest in New England and northern New York and generally in the glaciated portions of the areas studied.

#### SUMMARY OF RESULTS

The summary table that follows gives the results of measurements on topographic maps on a scale of 1:62,500. It includes about 22,000 areas covering 145,000 square miles. A total of 240,000 miles of stream length was measured, and nearly a million contours on the topographic maps were counted and translated into land and channel slopes.

Reference has already been made to general relationships between the topographic factors listed in the table. Each item is not necessarily unique, but it may reflect a condition that also influences the others, consequently other relationships between them may be found. For example, figure 52 shows that, in general, larger drainage areas are associated with flatter stream slopes; but average land slopes and

<sup>&</sup>lt;sup>13</sup> Grover, N. C., The floods of March 1936, pt. 1: U. S. Geol. Survey Water-Supply Paper 798, pp. 335-338, 1037

mean altitudes of drainage basins above outlets or gaging stations show a tendency to increase with drainage area. The points shown on figure 52 correspond to averages of groups of drainage basins within limited ranges in size. If individual basins were plotted on figure 52, material scattering of points would result, the basins plotting on the left being relatively flatter than those on the right. The average curve therefore provides a means for comparing the slopes and alti-

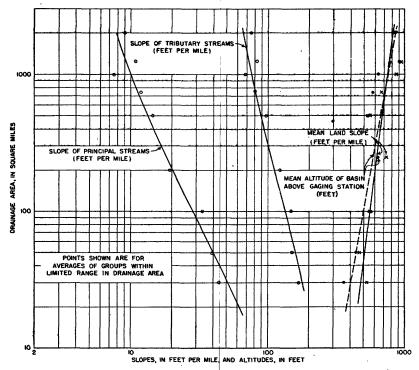


FIGURE 52.—Graph showing general variation in stream slopes and altitude in relation to size of drainage basin.

tudes of basins of different size. The divergent trends of the lines showing principal channel slope and average land slope indicate that the slope ratio of the basins analyzed tends to decrease with increase in drainage area.

A principal shortcoming of the computations of physical characteristics may be that it was not practicable to determine in detail the distribution of stream and land slopes and of lake and swamp areas within every area. Steep slopes on a few tributaries may increase the average slope considerably, yet these slopes may have little effect on flood-peak discharges. Moreover, the course of a river in a given

length may be characterized either by uniform slope or by a series of pools with intervening rapids or perhaps cataracts. The velocity in a stretch of uniform slope would probably be the greater if other factors were constant, as pools have a detention storage effect and the fall at rapids or cataracts imparts but little horizontal velocity to the water. A lake on the headwaters of a stream may have no noticeable effect, whereas a lake of the same size on the main channel near the lower reaches of an otherwise flashy stream may greatly modify flood discharges. Also, steep slopes or abnormally high elevations in the part of a drainage basin upstream from a lake may affect considerably the average land slope and the mean elevation of a basin, but the lake may decrease the flood discharges so much that the outflow from the lake would differ little from that of a basin in which the slopes and elevations were much less.

Storage capacity was not computed, as topographic maps furnish no information from which channel and lake cross sections at different stages can be determined, except that they might be crudely correlated with the stream slopes.

Summary of drainage basin topographic characteristics

	<b></b>		131(101	.100				1011011	.10		
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water (square	fatoT				0.08	1. 16	7. 24 1. 58 13. 42 3. 31	92. 32 35. 17 4. 62	8.32	8.6.2.58 4.88 5.5.5	12.00
rea of surfaces (so miles)	Squiswa				0	. 72	4.88 .58 1.32 2.57	25. 56 2. 56	5.38	4. 2. 2. 2. 2. 2. 2. 2. 2. 2. 2. 2. 2. 2.	10.58
Area surf mili	Lakes and reser- reser-	53.91	22.79 16.44 3.14	5.54 .17	60.	.44	2.36 . 06 . 72 . 74	83.84 9.61 1.96	2.94	3.93 .64 2.37 11.56	1.42
of land a hove a level)	muminiM	512 161	159 600 290	201 604 390	1,020	620	460 580 500 560 470	495 270 880	380 430	340 140 290 65	35
	меэМ	860 451	440 1, 765 1, 124	959 1, 760 935	2,804	2, 492	1,853 1,744 1,584 1,258	756 1,038 1,542	1, 101	820 868 868 268	126
Altitude (feet mean s	mumizsM	2,000	1, 463 4, 168 4, 237	4, 116 3, 535 2, 420	5, 249	5, 249	5, 249 4, 810 4, 810 2, 280 2, 920	2, 979 3, 165 2, 496	2, 722 2, 937	2, 378 2, 280 1, 990 755	400
-nirq smsər	to dtgned ( tta lagio (miles)	82. 2 28. 6	51. 1 73. 3 51. 6	62 20.8 18.4	18.7	23.9	52.4 13.0 29.8 15.3	39.9 91.3	31.	24.7 31.9 24.9 64	21.5
s)	Longest wat (miles	75. 0 30. 9	41.5 81.3 41.7	61.3 17.9 15.3	15.3	23.6	41.8 14.7 22.5 15.0 22.0	34.0 72.0 23.7	23.9 33.1	28.6 20.6 51.6	27.5
Channel slope (feet per mile)	Principal	2. 56	5.55 12.88 26.71	11. 95 79. 57 34. 5	117	76.2	21. 4 109. 2 54. 9 11. 1 33. 7	9.78 24.9	24.3 16.6	2.5 42.5 2.8 8.5	2.33
Channel slope (feet per mile)	Tributary	52. 7 51. 1	49 146 165	147 454 258	498	426	223 482 326 212 242	96 90.2	163 178	109 127 115 24	15.9
(feet	A verage	449 298	314 694 745	1, 044 818	1,598	1, 478	1, 161 1, 004 878 603 764	460 575 610	652 674	511 497 416 261	194
Land slope (feet per mile)	.8N	445 294	303 742 707	662 1,014 730	1,468	1,363	1,089 870 794 565 697	438 522 542	590 627	505 475 372 244	192
Land	E'-M'	456 304	322 630 797	1, 088 1, 886 867	1,774	1, 636	1, 262 1, 188 998 654 852	491 646 711	734	519 525 474 285	197
(səlin	ı əiduə) la Z	32, 073 2, 531	7,010 35,524 7,114	16, 641 1, 051 622	824	2,355	13, 156 412 1, 735 1, 081	7, 274 28, 221 669	1,970	2, 270 3, 154 1, 168 12, 390	1,714
ytisne	Stream de region) (Alim	1.18	1.08 1.42 1.79	1, 47 1, 57 2, 34	1,45	1.50	1.62 1.84 1.62 1.42 2.05	1.46 1.89 2.32	1.54 1.65	1.59 1.59 1.59	1.23
area iles)	əzaniar U m ərsups)	867 148	299 872 351	514 95 76. 2	104	192, 6	622 58. 8 142 57. 6 85. 8	363 766 54. 8	146. 2 129	157 170 107 405	124
	Name of gaging station	Fish River near Fort Kent, Maine.	Avaine. Passadumkeag River at Lowell, Maine. Dead River at The Forks, Maine. Carrabassett River near North Anson, Maine.	∞ ∞ <u>⊢</u>	East Branch of Pemigewasset River near	Pemigewasset River at Woodstock, N. H.	Pemigewasset River at Plymouth, N. H. Bakers River at Wentworth, N. H. Bakers River near Rumney, N. H. Squam River at Ashland, N. H. Smith River near Bristol, N. H.	Lake Winnipesaukee outlet at Lakeport, N. H. Contoocook River at Penacook, N. H. North, Branch, of Contoocook River near		Suncook River at North Chichester, N. H. Soulbegan River at Merrinack. N. H. North Nashua River near Leominster, Mass. Concord River near Leominster, Mass.	at Lowell, Mass. Ipswich River near Ipswich, Mass
	N O	1-7	1-34 1-52 1-54	1-58	1-153	1-153A	1-155. 1-169. 1-170. 1-171A.	1-175 1-183 1-186	1-188A	1-192 1-198 1-202 1-210	1-213

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Summary of drainage basin

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water (square	Total	1.02 6.42 12.36 1.06 26.19	1.17 1.39 5.98 15.25 1.50	1.44 4.52 12.90 1.88 7.54	1.00 4.52 4.10 13.16	25. 47 1. 18 2. 40 2. 44	3. 11 12 2. 83
of v faces (sc les)	sqmsw8	0.48 3.54 8.50 12.24	25.03 0.29 0.29	0 1.50 8.02 3.42	. 3. 10 1. 13 1. 1	9.28 0.76 3.16 42	2. 28 . 08 1. 33
Area of surfaces miles)	Lakes and reser- voirs	2. 2. 2. 5. 2. 2. 2. 5. 2. 0. 2. 5. 2. 0. 2. 5. 3. 0. 5. 5. 4. 0. 5. 5. 5. 0. 5. 5. 5. 5. 0. 5. 5. 5. 5. 5. 5. 5. 5. 5. 5. 5. 5. 5.	. 92 1.36 5.90 12.96 1.50	1.8.4.1.4 20.8.21 20.8.21	1.30 2.94 3.46 10.12	16.19 1.18 1.64 2.02	. 83 0 0 0 1. 50
f land bove level)	muminiM	19889 1988	70 480 270 130 360	360 200 15 15 160	280 240 240 240	220 110 110	1, 190 460 550 640 840
೦ ನೆ ನೆ	Mean	232 232 211 213 99	195 746 611 495 496	637 500 392 698 620	605 614 825 779 696	553 744 536 512 406	2, 512 1, 712 1, 801 1, 324 1, 402
Altitude (feet mean s	mumixaM	2880 2880 420 420	1, 400 1, 400 1, 400 1, 400	910 805 805 1, 280 1, 290	1, 015 1, 290 1, 280 1, 280 1, 280	1, 280 1, 080 940 860 760	9,2,3,8 9,2,8,8 9,2,3,00 9,2,00 1,00 1
prin- eams	Length of cipsl str (miles)	9.4 31.9 44.1 14.6 36.6	15.3 13. 26.5 48 10.5	11.4 19.9 39 27.9 49.6	16.9 33 17.8 24.7 45.3	69.5 13.1 18.3 20.6	22.3 52.4 32.4 20.2
ercourse (s)	tsw teagno.I gelim)	88.7 59.8 111.6	11.5 26.5 44.5 9.8	7.6 17.0 27.7 25.3 32.6	26.5 28.3 38.1	65.3 9.5 17.8 16.5	20.2 50.7 31.3 10.1
nnel (feet aile)	Principal	12.8 3.57 19.2 3.28	11.6 41.5 12.5 12.3 20.0	46.1 21.6 15.1 25.1 13.3	20.7 22.0 8.97 12.2 13.7	12.1 26.2 24.3 24.3	75.8 35.9 21.8 119.0 39.2
Channel slope (feet per mile)	Tributary	20.2 20.6 35.7 13.5 5	24.0 77.9 58.7 41.5	97.6 59.0 55.6 91.4 61.7	114 79.9 60.2 56.1 51.0	42.1 76.6 70.9 68.7 64.0	575 274 350 375 234
(feet	Ачетаве	301 236 176 176	371 337 337 307 301	283 260 378 344	334 341 436 415 369	328 348 285 305	1, 130 931 1, 295 890 588,
Land slope per mile)	.8N	257 210 223 159 105	129 308 279 262 248	275 213 228 312 293	303 330 319 291	261 274 237 220 271	1, 056 855 1, 236 693 486
Land	E-W.	359 270 273 200 129	158 458 414 370 374	293 323 298 471 414	377 424 579 549 478	420 449 355 319 351	1, 225 1, 034 1, 378 1, 168
(səlin	z si (cubic z	3, 638 8, 126 202 3, 738	327 164 1,861 8,184 139	3,416 11,775 6,815	2, 320 2, 320 2, 362 6, 332	22, 477 141 617 723 738	968 8, 575 4, 272 172 622
Square	Stream de (miles per mile)	2.34 1.27 1.37 1.27 1.28	1. 20 1. 69 1. 36 1. 32 1. 57	1.14 1.47 1.51 1.69 1.62	1. 61 1. 63 1. 46 1. 57 1. 55	1.56 1.62 • 1.61 1.51 1.51	1. 72 2. 09 1. 99
a 9 1 a iles)	eganiar U m eraupa)	23. 3 183 251 35. 2	42. 2 31. 3 139. 2 417 26. 1	27.8 93.3 199.8 121 401	76.2 169 93.8 157 331	711 27.7 58.7 83.5 88.6	88. 5 241 30. 5 80. 5
	Name of gaging station	Aberjona River at Winchester, Mass. Charles River at Charles River Village, Mass. Charles River at Walthan, Mass. Neponser River at Norwood, Mass. Taunton River at State Farm, Mass.	Wading River near Norton, Mass. Blackstone River at Worcester, Mass. Blackstone River at Worthbridge, Mass. Blackstone River at Woonsocket, R. I. Quinsigamond River at North Grafton, Mass.	Mumford River at East Douglas, Mass.  Branch River at Forestdale, R. I. Pawktrsef River at Cransfoon, R. I. Willimantic River near South Coventry, Conn. Shetucket River near Willimantic, Coun.	Hop River near Columbia, Conn. Natchaug River at Willimantic, Conn. Quincbaug River at Westville, Mass. Quincbaug River at Quincbaug, Conn. Quincbaug River at Putnam, Conn.	Quinebaug River at Jewett City, Conn. Little River at Buffumville, Mass. Five-Mile River at Killingly, Com. Moosup River at Moosup, Com. Yantic River at Yantic, Conn.	Ammonoosue River at Bethlehem Junction, N. H. Ammonoosue River near Bath, N. H. White River near Bethal, Vt. Ayers Brook at Randolph, Vt. Mascoma River at West Canaan, N. H.
	No.	1-213.5 1-218. 1-218. 1-220A	1-23 1-225 1-227 1-230	1–232B 1–237A 1–249A 1–264	1–272 1–275 1–279A 1–282 1–284	1-289 1-294A 1-295. 5 1-298 1-301	1-329A 1-332 1-334 1-336.8A 1-339A

822 822 822 823 8. 13. 28 13. 38 20. 32 1.69 2.69 1.18 38 1.38 1.41 3.49 3.49 3.09 1.02 <del>3</del>8888 2.13 នុ ~ :: · -:05.03 4,00 .4.0.1 8.03 8.03 8.03 8.03 .0 448284 9.88 2.388 2.388 2.388 2.388 **8%89** 1.01-160 ed ici ici ങ്ങ 285 2865 2865 2865 246 318 318 318 1,655 1,517 1,204 1,435 1, 110 1, 070 988 1, 038 1, 108 1,557 1,438 1,438 870 870 990 865888 1, 466 1, 195 1, 192 544 166 240 830 830 85 35 88 88 35 88 22.25.55 w, 8, r<sub>c</sub> c4 45-19-19.8 32.6 37.9 9.8 -10 00 00 00 82.25.82.83 5.55 ğ 8.7.4 45. 82.24.7.3 82.242.29 25.23 4883 প্ত co co co co 050-0 1-10404 10 00 A -07-CO 00 10 53.6 25.85.53 17.22.23 11.889.11 128,827 52.5 ଷ୍ଟ প্ল 17. 9 16. 4 17. 4 24. 1 24. 1 37. 2 61. 3 16.0 12.3 17.4 11.7 සුපසු 175 193 85 G 38.83 282 89 89 090 615 932 610 808 313 580 614 660 3310 328 328 314 468 171 186 186 567 577 508 702 318 347 301 151 282 278 278 278 87 87 545 530 628 232 232 766 884 562 211 733 666 751 952 274 125 446 919 025 109 109 677 998 413 216 -14,000 4.00.00 ಡ್ಪಪ್ಪ-21 32 2.17 1.92 1.72 1.94 \$18 36 36 36 71.2 420 41.8 £0 40 € 830 19 19 . 12. 362 සිනිනීසුව **\$25** Si 8,8,8,5 land, Vt. Millers River near Winchendon, Mass......
Millers River at South Royalston, Mass.....
Millers River at Erving, Mass.....
Sip Pond Brook near Winchendon, Mass....
Priest Brook near Winchendon, Mass.... East Branch of Tully River near Athol, Mass.
Moss Brook at Wendell Depot, Mass.
Deerfield River at Charlemout, Mass.
Deerfield River, excluding Somerset Reservoir, at Charlemout, Mass.
Deerfield River, excluding Farriman Reservoir, yat Charlemout, Mass. Ware River at Gibbs Crossing, Mass.
Chicopee River at Bircham Bend, Mass.
Weith River at West Ware, Mass.
Quaboag River at West Brimfeld, Mass.
Mill River at Springfield, Mass. Westfield Little River near Westfield, Mass.... Scantic River at Broad Brook, Conn... Farmington River near New Boston, Mass..... Farmington River, excluding Otis Reservoir, near New Boston, Mass. ------------------Ashuelot River near Marlboro, West River at Newfane, Vt.
Ashuelot River near Glismr, N. H.
Ashuelot River at Hinsdale, N. H.
Otter Brook near Keene, N. H.
South Branch Ashnelot River near M.
N. H. River near East Hartrord, Mascoma River at Mascoma, N. H. Ottauquechee River at North Hartl Sugar River at West Claremout, N. Black River at North Springfield, Waxtons River near Saxtons River, N Hockanum ton, 1-376A 1-380A 1-380.4A 1-383.2 1-384 1-387 1-389 1-390 1-394 1-396.8 1-399 

Summary of drainage basin topographic characteristics—Continued

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Apich	o redmuN otni asena otni ased aivided	51 52 72	69	51 50 50 50 50	52 74 33 53 49	234	45 63 105	53	49	39,49	27 51	34
	IstoT	0.91 1.03 70 5.10	10.96	26.29 3.77 1.98 2.88	1.81 1.81 1.24				1.47	. 74	0	85.
rea of water surfaces (square miles)	Sqmswg	0 .03 01.	1.72	9.42 2.26 1.02 0.62	0.12			-	1.27	0	0	0.16
Area surf mil	Lakes and reser-	0.91 1.00 4.74	9.24	16.87 1.51 2.26 37	. 19 1. 69 . 10 . 14 5. 85	40.97	3,84 18,55 13,84	8	. 20	74	0	288
land oove level)	muminiM	90 40 1,030 720	555	320 210 300 310	410 160 420 30 1,550	290	1, 544 700 905	1,248	620	374 790	370	870 530
Altitude of land (feet above mean sea level)	пвэМ	500 300 1,678 1,432	1, 264	981 849 568 1,029 685	1,016 801 870 459 2,193	1,848	2, 231 1, 450 1, 888	1,896	1, 700	778	1,650	1,835 1,649
Altitude (feet mean s	mumixsM	920 1, 900 2, 300 2, 660	2,660	2, 660 1, 737 1, 070 1, 680 1, 140	1, 700 1, 200 1, 660 5, 344	5,344	3,865 4,842 3,595	3, 595	3,816	3,020	3, 300	3, 143 3, 764
-ninq smsən	dagrad. 1 s agio 1 (selim)	21. 4 22. 3 14. 5 36. 7	55.6	99.3 34.6 17.6 33.1 19.5	21. 1 31. 4 11 26. 4 37. 1	111.7	36.3 44.5 53.9	23. 1	23.3	21.8	11. 5 53. 3	12.2 21.2
s) percourse	Longest was	17.5 23.2 12.2 43.0	68.7	111.2 33.0 19.4 34.7 17.1	21.8 40.4 10.1 19.7 24.4	93. 7	35.0 47.3 35.9	22.3	22.2	19.9	46.0	9.5
Channel slope (feet per mile)	Principal	40. 2 9. 42 78. 6 13. 1	8.45	10 17.1 17.1 31.4 33.4	41 17.5 72.3 32.2 13.0	12.37	32. 61 16. 54 35. 7	42.5	52.7	24.36	113	146 80.2
Cha Slope per 1	Tributary	74. 6 91. 2 165 135	79.8	54. 2 113 85. 2 88 153	175 103 147 92. 6 250	127	159 118 124	183	314	163 397	414	479 271
(feet	A verage	349 351 521 553	543	557 537 487 465 474	546 483 424 474 1,091	920	800 1,040 987	1,157	953	432 673	900	841
Land slope per mile)	.8N	286 274 486 513	487	477 480 407 388 366	475 399 334 368 1,010	882	883 998 961	1,074	827	390 544	533 795	719 653
Land	е-м.	433 452 567 607	621	666 736 593 574 624	644 598 552 614 1, 148	973	084 1, 100 1, 022	1,275	1, 120	494 842	1,047	1,005
(səlim	z sidus) ls Z	821 1,323 405 6,841	21,092	81, 138 3, 497 820 2, 431 710	4, 483 121 710 2, 282	89, 931	2, 503 13, 054 9, 640	1,258	1,552	894 268	142 11, 036	210 1,067
Tiene Square	ob maest2 reg selim) (elim	1. 62 1. 66 1. 91 1. 52	1.35	1.57 1.34 1.81 1.98 2.29	22.21 2.25 2.15 2.15 2.21 2.21	1.68	1.77 1.58 1.60	1, 37	1,51	1.68	1.79	1.92
8 1 8 g. lies)	oganiar C m oranar)	104. 7 109 57. 1 280	632	1, 545 204 68. 5 133 75. 3	71.9 245.8 24.0 77.5	1,664	160 527 491	114	152	90	$\begin{array}{c} 31.4 \\ 510 \end{array}$	39.1 111
	Name of gaging station	Salmon River near East Hampton, Conn Quinnipiae River at Wallingford, Conn Housatonic River at Coltaville, Mass Housatonic River near Great Barrington,	Housatonic River at Falls Village, Conn.	Housatonic River at Stevenson, Conn. Tenmile River near Gaylordsville, Conn. Shepaug River near Ronsville, Conn. Shepaug River near Roxbury, Conn.	Naugatuck River near Thomaston, Conn Naugatuck River near Naugatuck, Corn Leadlinie Brook near Thomaston, Conn Saugatuck River near Westport, Com Hudson River near Newcomb, N. Y	Hudson River at Hadley, N. Y.		3 (	Batten Kill at Arlington, Vt.	ms, Mass	Adams, Mass  Hoosic River near Eagle Bridge, N. Y	North Branch of Hoosic River at North Adams, Mass. Walloomsac River near North Bennington, Vt.
	No.	1-432 1-435 1-444	1-446	1-450 1-453 1-455 1-459	1-462 1-463 1-464 1-466	1-475	1-485 1-488 1-409	1	1-496	1-501	1-506	1-508

78 66 61	61. 51. 65 69	49 118 49	49 57	66 49 56 56	45	148 33 7 60	22	853 853 80 80 80 80 80 80 80 80 80 80 80 80 80	56 49 31 25 55	53 70 119 65	338
				3.00		0			1.82	9.25	
				2.55		0		0	1.81	9.24	_
12.07 4.35 1.52	. 10 . 65 1. 11 3. 09 1. 06	. 08 13.90 3.29	$\frac{.27}{2.13}$	2.78 .45 .20 .79	8.	. 99 0 0 82.	0.16	12285	.01 .20 .18 .18	. 56 0 2.30	33.08
857 850 84	648 55 390 196 140	208 360	265	8828	130	29 113 275 192	52	294 30 10 10	22 8 00 10 10 10 10 10 10 10 10 10 10 10 10 1	55 40 955	420
1, 131 1, 821 897	1, 347 930 695 624 493	319 456 822	534	654 275 204 778	292	370 229 480 587	378	732 154 137 265 113	112 106 84 64 93	125 108 79 1,849	1,555
3, 626 2, 780 2, 653	2,885 3,863 1,809 2,273 1,440	857 1, 496 1, 406	1,092	1, 420 710 643 1, 227	1, 220	1, 220 680 660 1, 122	1,160	1, 205 540 563 563 561 391	308 346 1190 205	199 177 131 3, 905	3, 905
112.3 38.3 57.1	17 46.3 26.7 74.7 34.2	17.7 92.9 27.2	12.1 25.1	30.4 22.4 19 23	32. 7	60.3 10.3 1.6 13.4	37.7	14.9 36.2 36.4 6.7 18.9	11.1 18.7 15.8 15.8	19 22.7 7.4 75.1	173.4
28.1 44.4	15. <b>9</b> 46. 7 23. 9 72. 6 30. 6	11.9 59.9 31.0	11.3	35.4 18.0 18.4 23.7	35.1	58.2 7.8 11.6	33.3	17.6 26.6 35.6 7.47 5 9.8	12.3 21.8 14.7 17.1	16.7 16.9 7.3 62.3	157.9
11.04 51.8 26.8	56.4 26.8 11.2 4.7 16.1	9.8 7.1 14.9	83.5 12.3	15.0 19.2 19.9 31.3	19.6	10.1 24.1 73.7	23.80	28.46 5.80 4.15 47.3 13.6 5	8.02 6.99 8.98 5.57 6.77	6.05 5.46 7.2 13.3	7.66
73 95. 7	309 144 102 63. 7 70. 2	67.1 49.6 88.9	181 21.8	112 66.9 114 151	. 801	71.6 92.3 300 191	75.7	125 23.3 29.3 145 46.4	27.1 13.6 18.2 10. 15.6	12.8 10.1 11.7 126	90.1
474 543 679	770 702 554 449 471	294 533 602	574	8661 309 501	451	378 259 473	437	443 140 158 373 244	121 158 71 70 70	57 60 60 92	277
497 568 612	820 660 500 412 358	315 473 578	562 648	596 154 305 497	453	370 262 529	427	450 154 167 379 250	132 151 72 57 69	53 61 982	210
441 526 765	709 748 636 498 663	266 609 632	580 860	825 320 506	449	389 255 414	453	434 131 152 365 240	114 166 70 54 71	53 61 59 1, 241	860
- 65, 630 3, 475 8, 756	832 9, 603 1, 769 26, 284 3, 335	365 25, 065 1, 648	1, 307	2, 927 600 437 886	2, 408	12, 849 109 1, 84 163	2, 911	2, 128 4, 288 4, 288 40 275	318 1, 260 474 734 693	1, 036 89 26, 360	244, 482
1.95 1.75 1.26	1.47 1.38 1.97 1.60 1.35	1.89 1.73 1.90	1. 70 1. 76	1.67 2.25 2.02 1.55	1.61	1.85 1.86 1.43 2.37	2.00	1.98 1.58 1.72 1.08 2.41	1.98 1.22 1.39 1.00	. 96 1.14 1.16 1.38	1.37
1, 348 261 329	98 386 144 711 182	55.4 785 116	29.4	160 54. 6 . 40. 9 65. 3	147	25.7 25.7 26.2	190	32. 8 171 258 9. 75 48. 5	43.4 124 56 70.5 64	56.3 113 22.3 783	3,076
Mohawk River near Little Falls, N. Y. Esst Canada Creek at Dolgeville, N. Y. Kinderhook Creek at Rossman, N. Y.	Catskill Creek at Oak Hill, N. Y. Rondout Creek at Rosendale, N. Y. Wallkill River near Unioville, N. Y. Wallkill River at Gardiner, N. Y. Wappinger Creek near Wappinger Falls, N. Y.		Whippany River at Morristown, N. J Ramapo River near Mahwah, N. J		N.J. South Branch of Raritan River at Stanton, N.J.		Hills, N. J. North Branch of Raritan River at Milltown,		Manasquan River at Squankum, N. J. Toms River near Toms River, N. J. Cedar Creek at Lanoka Harbor, N. J. Batsto River at Batsto, N. J. East Branch of Wading River at Harrisville, N. J.	Great Egg River at Folsom, N. J. Mandrice River at Norma, N. J. Manantico Creek near Millville, N. J. Bast, Branch of Delaware River at Fishs	Eduy, N. I. Delaware River at Port Jervis, N. Y
1-515 1-536 1-549	1-550 1-562 1-563.5 1-566 1-569	1-577 1-582 1-583.5	1-587	1-595 1-606 1-609 1-611	1-614	1-616 1-619 1-619.5 1-620	1-622	1-623 1-624 1-625 1-625:5	1-634 1-634 1-636 1-636	1-639 1-640 1-641	1-646

Summary of drainage basin topographic characteristics—Continued

	wpich	to 19dmuM otalizesie gw nized bebivib	52 40	114	84	ខ	15 25 25 4 25 4	12 43 44 12 44	57 47	119 48	73 47 39	51
	water (square	Total										
	of faces es)	Syamps										
	Area surfac miles)	bns seas. Teser- Toriov	36.30	38.	16.21	3.55	2. 0.0 2. 32 558 558 558	. 16 4. 75 5. 08	.30	0.75	0.00 10.04	.02
	land oove level)	muminiM	1,335	955	874	874	645 348 410 335 560	404 296 1,060 210 260	848	798	128882	150 85 85
	ltitude of land (feet above mean sea level)	Mean	1,911	1,738	1,429	1,358	1,644 819 883 666 706	689 539 1, 724 1, 166 463	97 275 97	873 1, 293	438 113 305 321 86	447 90 340.
	Altitude (feet mean s	mumixsM	3, 253	3, 365	2,654	2, 654	4, 204 1, 653 1, 600 1, 600 1, 127	1, 248 1, 141 2, 320 2, 320 960	340 660 213	2, 020 2, 020	1,200 172 595 600 166	1, 100 164 640
	prin-	Length of cipal str (miles)	201.7	09	54.3	41.1	23.6 25.3 27.3	27.6 13.7 48.3 72.1 21.4	88.8 86.8 8.8	95 16.9	49.3 15.8 13.4 7.5	41 7.2 20.1
9		Longest was	209. 2 25. 8	73.3	34.6	34.6	8,22,23 8,22,38 8,73 8,73 8,73	28. 4 13. 5 41. 5 99. 3	20. 5 37. 5 18. 4	79.8	27.8 7.8 15.7 8.2	38.7 7.2 20.8
-	inel (feet	Principal	16.40	82.8	18.7	19.1	30.7 31.9 40.67 11.9 10.75	7.07 15.55 22.9 16.5 12.9	4.75 8.0 6.5	6.77	14. 08 20. 4 23. 1 24. 6 13. 73	12.8 11.1 20.1
	Channel slope (feet per mile)	Tributary	154	130	75.4	83	129 174 154 154 82 78.1	80.3 122 59.4 55.3 40.5	22 43.7 12.7	70. 7 202	65.2 54.2 104 106 47	51.5 39.3 96.3
	(feet	<b>А</b> чега <b>ge</b>	716	905	469	498	633 617 571 469 515	497 526 283 497 233	109 225 62	605 621	360 98 484 412 148	434 144 389
	Land slope (feet per mile)	.8N	823	857	426	422	705 593 630 467 476	476 535 277 557 222	118 237 66	654 780	364 98 463 438 155	459 154 373
	Land	EW.	1,005	964	527	601	848 641 502 472 534	509 515 288 416 244	99 212 57	540 443	357 96 509 378 138	402 130 414
	(səlim	ı əlduə) la Z	610, 833	20, 781	9, 260	4,857	5, 520 827 747 2, 093 138	1, 611 304 7, 183 57, 905 1, 263	810 4, 077 1, 197	49, 456 256	4, 277 12 335 267 76	5, 283 63 604
•	square	Stream de (miles per mile)	1.43	1.38	1.27	1.38	1.41 1.58 1.80 1.59	1.70 1.78 1.25 1.25 1.75	1.68 1.18	2.07 1.36	1.68 2.19 1.73 1.73	2.26 1.75 3.03
	a e a a iles)	eganiar U m eraups)	6, 344	593	519	280	222 65.1 64.4 126 31.4	108 36. 2 322 . 1, 280 97. 4	89.4 210 111	1, 147 42. 9	279 6. 75 33. 3 81. 9 19. 3	287 14.6 52.6
		Name of gaging station	Delaware River at Riegelsville, N. J. West Branch of Delaware River at Delhi,	West Branch of Delaware River at Hale Eddy,	Lackawaxen River at Hawley, Pa., (including	Wallenpaupack Keservorr). Lackawaxen River at Hawley, Pa., (excluding Wallenpaupack Reservoir).	Neversink River at Oakland Valley, N. Y. Flat Brook near Flatbrookville, N. J. McMidnaals Creek at Stroudsburg, Fa. Paulins Kill at Blarstown, N. J. Pequest River at Huntsville, N. J.	Pequest River at Pequest, N. J. Beaver Brook near Belvidere, N. J. Lehigh River at Tannery, Pa. Lehigh River at Bethehem, Pa. Tohickon Creek near Pipersville, Pa.	Assurphik Creek at Trenton, N. J. Neshaminy Greek near Langhorne, Pa. North Branch of Rancocas Greek at Pember-	Schuylkill River at Pottstown, Pa Little Schuylkill River at Tamaqua, Pa	Perktomen Creek at Graters Ford, Pa. Mantua Creek near Pitman, N. J. Grun Creek at Woodlyn, Pa. Ridley Creek at Moylan, Pa. Oldmans Creek near Woodstown, N. J.	Brandywine Creek at Chadds Ford, Pa Salam Creek at Woodstown, N. J. Big Elk Creek at Elk Mills, Md
•		No.	1-648 1-657	1-658	1-664.5		1-668 -671 -672 1-673	1-674 1-675 1-677 1-677	1-683 1-685 1-686	1-6881-691	1-693 1-694A 1-696 1-696	1-701 1-702A 1-706

97 198	67 49 68 128 61	51 210 24 41	45 47 47 47	64	51 52	57 57	75 49 57 57	82	55 67 47 165	15 65 51	30	22
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10.63	23.01 3.11 6.28 2.61		2. 52 2. 53 . 37 . 08	.05	0.13	0.15	8.08.	82	000	000	0	.85
1, 152	1,175 999 880 920	970 934 1, 190	1,060 1,010 775 560 1,225	582	420 820	754 620	792 424 552 354 260	970	1,005 651 1,022 239	970 270 155	270	_
1,489	1, 508 1, 430 1, 315 1, 430	1, 280 1, 573 1, 579 1, 770	1, 597 1, 520 1, 307 1, 189 1, 589	1, 241	1,488	1,477	1, 487 985 959 831 507	2, 455	1, 932 1, 633 1, 474 623	1,366 496 349	384	_
2, 301 2, 740	2, 080 2, 120 2, 160 2, 160	2, 540 2, 540 2, 540 2, 240	2,2,2,2,2,2,2,2,2,2,2,2,2,2,2,2,2,2,2,	2, 440	1, 813 3, 136	2, 620 2, 400	3, 136 1, 240 1, 320 1, 320	4, 150	2, 9, 022 1, 942 1, 980	1,880 880 600	260	
61.1	197 22. 71. 7 112. 3 79. 4	40.4 69.5 122 11.5	25.2 43.7 17.4 48.6 50.1	62.3	42. 5 32. 4	23.65 23.03	85.3 10.3 41.3	56.6	13.4 37.8 10.4 57.9	3.2 17 21.6	8	76.6
128. 4.51	176.9 25.0 58.3 76.5 58.8	31.7 44.4 57.3 93.0	19.2 42.2 18.4 30.6 39.3	43.4	36.8 31.9	28.1	78.6 40.4 11.2 80.8 35.7	48.7	17.8 33.7 11.5 63.4	3.5 18.6 21.7	6.9	83.2
3.90	15.40 5.16 4.63 6.99	17.45 12.30 7.49 6.1	36 10.30 12.30 13.10	16.34	16. 42 16. 80	24. 2 19. 5	7.7.1 17.2 6.3 8.3	55.2	55.7 51.7 39.0 5.1	223 14.9 15.2	22. 2	_
94.9 64.9	93.5 75.3 59.1	134 78. 5 101 50. 2 242	141 80.1 152 148 78.5	=======================================	89.5 143	218 167	88.2 136 298 71.6 54.5	198	337 192 409 69. 2	356 78 73. 4	80.2	_
621	576 585 559 572	641 782 780 646	685 603 442 778 710	268	758 873	900	895 1,084 465 363	746	850 1,053 1,072 438	715 524 410	315	
566	513 524 489 481	575 793 783 688	716 545 482 802 744	954	958 835	902	1,032 1,284 1,284 486 403	744	775 999 891 417	617 483 418	278	_
671 710	626 661 653 693	740 768 776 590	645 683 394 755 664	829	494 921	897 902	997 811 872 449 337	748	950 1, 125 1, 329 464	808 567 399	361	1
8, 829 163, 292	670, 664 1, 405 15, 249 57, 740 23, 148	2, 961 19, 033 43, 398 140, 000	1, 155 9, 429 4, 648 6, 191	12, 376	3, 283 5, 340	3, 180	35, 640 3, 971 19, 025 6, 368	7,717	4, 107 4, 107 165 26, 587	10 901 781	82	70,663
1.54	1.59	1.36	1. 63 1. 47 1. 67 1. 80 1. 72	1.34	1.55	1.50	1.86 1.89 1.25 1.25	1.36	1.72 1.49 1.67 2.02	2.01 1.80 1.92	1.56	_
351 2, 240	7, 797 103 518 1, 492 1, 735	186 770 1,370 2,530 30.5	114 472 45.8 274 315	559	162 291	220 128	756 200 21. 6 470 322	287	72. 4 247 30. 2 817	5.7 82.3 62.2	21.3	1,690
Susquehanna River at Colliersville, N. Y	Susquehanna River at Towanda, Pa. Oaks Creek at Index, N. Y. Unadilla River at Rockdale, N. Y. Chenango River near Chenango Forks, N. Y. Tioughnioga River at Itaska, N. Y.	Owego Creek near Owego, N. Y. Tioga River at Lindley, N. Y. Tioga River near Erwins, N. Y. Chemung River at Chemung, N. Y. Canisteo River at Arkport, N. Y.	Tuscarora Creek near South Addison, N. Y Coloticton River near Campbell, N. Y Wapwallopen Creek near Wapwallopen, Pa Fishing Creek near Bloomshurg, Pa West Branch of Susquehanna River at Bower, Pa.	North Bald Eagle Creek at Beech Creek Sta-	Mahantango Creek East near Dalmatia, Pa. Frankstown Branch of Juniata River at Williamshure Pa	Little Juniata River at Spruce Creek Pa Standing Stone Creek near Huntingdon, Pa	Raystown Branch of Juniata River at Saxton, Pa. Sherman Creek at Shermandale, Pa. Clark Creek near Carsonville, Pa. Conodoquinet Creek near Hogestown, Pa	North Branch of Potomac River at Blooming-	Georges Creek at Franklin, Md Wills Creek near Cumberland, Md Evitts Creek near Bedford Valley, Pa Monosacy River at Jug Bridge, near Frederick, Md	Owens Creek at Lantz, Md	Northwest Branch of Anacostia River near	t Red House, N. Y
1-707	1-712 1-717 1-723 1-727	1-740 1-742 1-743 1-746	1-758 1-761 1-777 1-778.5	1-790	1-798	1-802.5	1-818. 1-818. 1-818.5. 1-820.	1-851	1–868 1–866 1–867 1–903	1-904 1-905 1-910	1-913	3-2

Summary of drainage basin topographic characteristics—Continued

wpich	o nednuví sres into saw nizad bebivib	123 175 52 57 49	100 51 98 53 99	280 53 49 73 59	122 35 136 63 50	49 106 275 106 47	51 75 83 13	65 59
water square	IstoT				2.3		4.11	
of aces (ass)	squisws				2.3		2.18	
Area c surfac miles)	Lakes and reser- voirs	25. 53 25. 92 . 16 . 08	2.71 .03 0 1.71	2.05 .01 .19 0	0.05	0.03	1.93 1.93 .02 0	.03
of land a bo ve	muminiM	1, 195 1, 220 1, 100	1, 021 1, 220 1, 100 1, 080	950 920 972 928	3,060 3,060 1,670 2,175	975 828 752 760 860	805 940 952 972 870	952 850
	Меап	1, 595 1, 619 1, 385	1, 307 1, 491 1, 458 1, 220 1, 979	1, 590 1, 535 1, 477 1, 950 1, 222	1, 209 3, 370 2, 138 2, 247 2, 505	1, 110 1, 038 1, 051 1, 078 1, 138	1, 254 1, 055 1, 132 1, 120 1, 120	1, 072 1, 101
Altitude (feet mean s	mumixsM	1, 982 2, 140 1, 880	1, 880 2, 240 1, 700 2, 949	2, 949 2, 480 4, 775 1, 900	1, 960 4, 375 3, 340 3, 213 3, 027	1, 380 1, 380 1, 540 1, 470 1, 480	1, 600 1, 300 1, 400 1, 320 1, 140	1, 320
prin- sms91	lo dinad. Legio (eslim)	184.7 259.2 45.4 35 66.8	68.3 25.1 56.9 33.2	125. 5 62. 7 35. 6 115. 5 52. 8	75.5 27.4 107.5 42.8 9.2	30.8 91.2 145.1 69.8 23.4	20.02 20.02 20.03 20.03 20.03 20.03	21.7
ercourse s)	kw tesgno.I selim)	178. 4 214. 7 36. 9 25. 5 68. 9	95.9 22.1 27.2 27.2 50.4	106.3 39.0 36.2 113.7 66.5	86.7 19.8 92.2 50.3 8.5	30.4 82.2 111.1 58.8 20.9	25.3 25.2 25.2 1.8 1.8	20.9 50.8
nnel (feet aile)	Principal	11.2 16.8 6.6	3.6 15.6 7.7 8.2 29.1	9229 1299 1294 1299	3.3 13.3 20.3 16 73.59	6.3 3.1 3.4 2.87 10.5	11. 52 6. 21 6. 64 8. 66 54. 4	7.1
Channel slope (feet per mile)	Tributary	59.8 101 33.6	34.6 97.2 87.3 67.8	97.6 64.8 155 84.7	43.7 114 88.5 84.1 151	25.23.23 29.23.29 29.23.29	36. 4 39. 7 35. 3 34. 35. 3	54 34.3
(feet	Average	467 677 339	323 684 764 845 596	680 612 724 1, 202 1, 265	1, 277 520 705 553 583	211 158 211 197 290	214 233 233 922	191 345
Land slope (feet per mile)	.av	442 721 311	308 702 777 766 575	663 608 724 1, 179 1, 270	1, 290 500 695 527 495	199- 151- 208- 207- 306-	413 196 529 233 794	200 379
Land	E-W.	500 633 378	344 665 750 922 622	704 619 724 1, 231 1, 258	1, 260 541 719 586 649	224 166 216 190 269	480 225 615 1, 098	180 322
(səlin	ı oiduo) la Z	495, 220 800, 000 5, 626 3, 543 21, 800	48, 700 1, 700 8, 169 2, 700 20, 880	8, 230 8, 230 5, 695 71, 600 12, 930	29, 520 795 65, 400 8, 373 108	3, 882 36, 418 123, 609 18, 071 1, 026	10, 808 2, 297 3, 608 2, 115 1, 71	1, 293 11, 232
square	Stream de (miles per mile)	1. 38	1.43 1.78 1.90 2.31 1.65	1, 75 1, 70 1, 64 1, 80 2, 07	2.29 1.11 1.71 1.68	1. 52 1. 46 1. 48 1. 56 1. 56	1. 60 1. 56 2. 13 2. 03 3. 37	1.77
a s t a (zəli	əzaniar C m əranpe)	5, 982 7, 671 233 629	1, 028 153 321 191 715	1, 825 390 265 1, 340 384	750 86.2 1,326 24.5	247 899 2, 235 588 104	406 165 254 175 1.64	120
	Name of gaging station	Allegheny River at Franklin, Pa. Allegheny River at Parkers Landing, Pa. Brokenstrawe Creek at Youngsville, Pa. Tionesta Creek at Lynch, Pa. French Creek at Seggerstown, Pa.	French Creek at Utica, Pa.  Mahoning Creek at Punxsutawney, Pa.  Mahoning Creek near Dayton, Pa.  Crooked Creek at Idaho, Pa.  Conemaugh River at Seward, Pa.	Kiskiminetas River at Vandergrift, Pa. Blacklick Creek at Blacklick, Pa. Loyalhanna Creek at Ivew Alexandria, Pa. Tygart, River at Fetterman, W. Va. West Fork River at Clarksburg, W. Va.	West Fork River at Enterprise, W. Va. Blackwater River at Davis, W. Va. Youghfogeney River at Connellsville, Pa Casselman River at Markleton, Pa. Big Piney Run near Salisbury, Pa.	Mahoning River near Berlin Center, Ohio Mahoning River at Youngstown, Ohio Basver River at Wampum, Pa. Shenango River at Sharpsville, Pa Little Shenango River near Greenville, Pa	Slippery Rock Creek at Wurtemburg, Pa. Tuscarawas River at Clinton, Ohio. Sandy Creek at Waynesburg, Ohio. Nimishilen Creek at North Industry, Ohio. Home Creek near New Philadelphia, Ohio.	Jerome Fork at Jeromeville, Ohio
	N0.	3-3 3-4 3-18 3-18.5	3-26 3-33.5 3-34.5 3-42.5	3-47 3-55 3-59 3-63 3-70	3-71 3-81 3-95 3-98 3-100	3-111 3-114 3-117 3-125.5	3-135 3-141 3-148.5 3-150	3-158

£42 24 25	53 112 104	69 61 115 74 258	47 45 58 69 172	55 53 67 180	61 120 215 37 40	31 108 83 83 83	. 164 65 92 78 134	87 81 87 89
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796 785 745	754 702 730 730 638	1, 444 2, 080 1, 522 1, 995 675	2, 185 2, 076 1, 856 925 608	635 674 574 704 561	694 639 569 618 601	910 905 879 830 820	708 735 715 720 667	584 630 500 516 928
998 958 960	1,002 1,195 1,081 915 879	2, 392 2, 764 2, 764 2, 703	3, 166 3, 079 2, 761 2, 587 1, 856	1, 581 1, 370 1, 761 1, 723 1, 399	1, 665 1, 649 1, 378 1, 129 1, 111	997 972 1, 038 1, 018 1, 061	989 927 950 963 970	856 749 868 884 884 1,036
1, 280 1, 280 1, 260	1,2,800 1,240 1,240 1,240	4, 4, 300 4, 4, 4, 4, 4, 7, 7, 10	4, 710 4, 524 4, 372 4, 839 4, 839	3, 400 1, 350 1, 100 3, 536 3, 536	3, 095 3, 765 2, 500 2, 500	1,220 1,050 1,420 1,420 1,420	1, 420 1, 220 1, 300 1, 343	8 1, 200 4 1, 200 1, 200 1, 200 1 1, 550
46.2 4.8 4.7	62.55 62.55 64.4.65 64.4.60	63.3 75.7 128.1 51.2 116	29. 1 34. 8 47. 4 65. 6 142	61 50.9 85.5 85.2 119.7	63 94.3 137.7 57.9 64.1	33.8 27.4 64.4 33.8	87.5 86.5 84.8 84.8	34. 15. 101. 38.
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3.01 19.23 9.97	8. 21 6. 99 3. 97 2. 44	22. 76 17. 19 10. 52 27. 19 25. 15	44. 12 72. 76 18. 59 35. 41 11. 56	24.1 18.21 2.28 11.26 7.01	18.06 13 8.02 7.79 7.30	2. 19 4. 8 5. 4 12. 04	6. 47 6. 75 5. 28 7. 32 7. 51	7. 07 7. 06 6. 72 8. 5 2. 52
39.5 127 74.5	26.4 76.6 54.6 32.7 29.5	108 154 114 275 148	347 259 130 285 145	140 198 26. 5 130 118	171 113 86. 1 120 120	19.4 11.9 17.8 18.1 29	26.6 28.7 16.8 24.2	59.9 64.5 19.6 21.4
417 . 799 . 660	363 1, 642 1, 673 1, 673 742	1, 338 1, 575 1, 455 1, 335 1, 321	1, 318 1, 414 1, 183 1, 768 1, 770	2, 218 2, 244 2, 244 2, 002 2, 077	2, 563 2, 242 2, 325 2, 325 2, 339	102 60 110 140	131 146 114 96 227	1, 016 806 216 183 126
414 731 667	358 1, 610 1, 670 1, 670 746	1, 332 1, 606 1, 474 1, 349 1, 302	1, 373 1, 481 1, 133 1, 774 1, 722	2, 149 2, 169 853 1, 993 2, 061	2, 511 2, 268 2, 211 2, 077 2, 496	100 56 102 118 134	130 149 116 107 226	983 802 218 162 118
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12, 196 129 1, 891	16, 414 10, 822 35, 364 10, 582 40, 234	15, 130 17, 475 91, 988 5, 597 74, 848	2, 041 1, 569 7, 542 11, 004 88, 456	12, 726 4, 740 23, 203 32, 960 91, 646	15, 246 56, 421 149, 121 5, 436 6, 708	5, 373 9, 778 13, 733 3, 533	17, 090 5, 060 21, 339 9, 800 31, 960	6, 271 551 60, 323 16, 837 13, 509
1.87 2.22 2.26	1999999	2. 03 1. 12 1. 46 1. 56	1. 40 1. 26 1. 56 1. 56	1.87 2.12 2.10 2.10	2. 09 2. 17 1. 77 1. 64	1.39 1.16 1.37 1.37 1.42	1.32 1.37 1.62 1.05	2. 09 2. 17 1. 62 1. 73 1. 44
466 27.5 140	672 386 913 460 944	438 540 1, 357 236 1, 315	128 150 287 281 1, 145	393 267 587 762 1, 226	389 1, 225 2, 150 198 209	255 73.3 387 438 195	544 190 533 331 808	286 76.5 1, 195 477 545
Mill Creek near Coshocton, Ohio	Licking River at Toboso, Ohio.  Little Kanawha River at Glenville, W. Va.  Little Kanawha River at Grantsville, W. Va.  Hocking River at Enterprise, Ohio.  Hocking River at Athens, Ohio.	Bluestone River at Lilly, W. Va.  Greenbrier River at Buckeye, W. Va.  Greenbrier River at Alderson, W. Va.  Gauley River at Camden-Gauley, W. Va.  Gauley River above Belva, W. Va.	Williams River at Dyer, W. Va. Cherry River at Fenwick, W. Va. Meadow River at Nallen, W. Va. Elk River at Centralia, W. Va. Elk River at Queen Shoals, W. Va.	Coal River at Ashiord, W. Va. Little Coal River at Madison, W. Va. Raccoon Creek at Adamsville, Ohio. Guyandot River at Man, W. Va. Guyandot River at Branchland, W. Va.	Levisa Fork at Fishtrap, Ky Levisa Fork at Pikeville, Ky Levisa Fork at Pankville, Ky Johns Creek near Prestomburg, Ky Johns Creek near Van Lear, Ky	Scioto River at Larue, Ohio	Big Walnut Creek at Rees, Ohio	Sait Creek near Londonderry, Ohio Little Sait Creek near Jackson, Ohio Little Miami River at Milford, Ohio East Fork of Little Miami River at Perintown, Ohio Miami River at Sidney, Ohio
3-161 3-163 3-166	3-167 3-169 3-170 3-175	3-203 3-205 3-206 3-208 3-208	3-213 3-216 3-217 3-220 3-222	3-224 3-227 3-230 3-232 3-233	3-235.7 3-235.7 3-238.5 3-238.5 3-238.5	3-243. 3-250.5. 3-252. 3-253. 3-254.5.	3-255 3-256 3-257 3-257 3-258	3-261.5 3-262 3-267 3-269 3-278

Summary of drainage basin topographic characteristics—Continued

Apich	rotni 28918 28 w nised bebivib	76 67 91 51	29	74 38 53 131	29 33 32 19	56 112 220	69 56	71 230 55 39	140 53	63 8 11.8
water b	IstoT to redmuM		-					.		
of v	sqmsw8									
Area of surfaces miles)	Lakes and reser- voirs	0.03	9.	20.00	8000	888	88	7.85 40.1 01 76	18, 58 2, 15	3, 14 55, 71 11, 25
land o o ve evel)	muminiM	850 952 732 601	485	790 955 784 578	990 610 675 605 1,060	1, 435 655 553	800	490 930 375 1, 488 1, 195	286 736	361 1,500 226
ea l	фвэМ	1, 032 1, 069 1, 079 910 927	845	1, 089 880 880	1, 120 1, 055 986 1, 111 1, 349	1,877 1,741 1,527	1, 491 1, 319	1, 408 1, 494 1, 576 2, 076 1, 650	1,049	1, 272 2, 036 1, 311
Altitude (feet mean s	mumixeM	1, 220	1,060	1, 380 1, 340 1, 340	1, 280 1, 860 1, 920 1, 920	2,565 2,565 2,565	2, 140 2, 028	2, 280 2, 2, 8, 2, 700 2, 4, 670 3, 670	2,460 2,400	2, 620 4, 621 3, 305
prin- sms97	Length of cipal st (miles)	93.28 83.29 83.88	111.8	27.9 24.4 72.9 130.7	31.1 30.9 40.9 8.6	35 73.9 105.3	26.8 33.5	35.2 41.9 131.8 60.5 26.5	127.9 64.1	66. 2 72. 7 106. 3
	sw teagno.I	46.5 36.1 19.7 39.6	85.5	28.28 11.58.1 119.53	39.7 26.0 33.0 7.9	29.4 85.1 101.9	29.5 4	38.3 40.0 118.5 39.5 24.2	124.3	58.7 78.5 71.6
nnel (feet nile)	Principal	5.11 5.56 8.29 18.2 15.7	5.23	9.18 9.34 6.15 3.96	15.4 15.05 26.38 22.5 69.2	15.7 14.8 10.58	82.9 23.18	45.06 28.6 17.09 12.36 33.74	11. 29 24. 18	28.3 6.01 25.18
Channel slope (feet per mile)	Tributary	112 14.8 27.8 50.4 36.3	23.	23.23 24.7.1 12.33	19.6 78.3 72 112 229	107 78.3 98.2	166	56.1 61.6 56.9 73 61.2	39. 7 60. 6	47.7 81.9 48.8
(feet	Ачегаве	124 192 192 220	324	021 106 93	261 359 295 426 476	717 566 539	589 408	363 397 469 730 489	481	469 764 417
Land slope per mile)	.8N	103 136 193 221	328	165 85 103 91	262 340 294 409 412	736 562 526	514 374	322 388 482 482 870 519	470	450 819 447
Land	E-W	92 117 191 174 216	320	177 112 112 95	260 386 295 451 560	697 570 557	641 454	483 404 450 634 472	498	495 691 378
(səlim	z si (cubic	12, 496 4, 008 9, 232 714 6, 070	21, 27.1	1, 781 1, 346 9, 138 74, 555	230 3, 510 1, 222 2, 855 88	4,989 49,429 74,938	1,578 2,014	3, 497 5, 964 133, 714 8, 704 1, 052	68, 586 5, 176	10, 163 28, 408 24, 457
vaiene square	db mae112 19q selim) (elim	1.33	1.46	1.16 1.40 1.38 1.55	. 65 1. 71 1. 94 1. 64 1. 06	1.39 1.45 1.49	1.52	2, 26 1, 87 1, 80 2, 18 1, 68	1.62	1.59 1.88 1.51
	əşaniar C m əranpa)	502 - 195 485 70 315	463	113 89. 8 299 1, 248	42 145 93.3 136 22	309 1, 017 1, 419	153 124	189 295 1, 876 365 85	973 258	335 722 616
Name of gaging station		Stillwater River at Pleasant Hill, Ohio.  Greenville Creek near Bradrord, Ohio.  Mad River near Springfield, Ohio.  Wolf Creek at Jayton, Ohio.  Talawanda Creek near Sevenmile, Ohio.	Elkhorn Creek, at Knights Bridge, near Frank-	Cedar Creek near Cedarburg, Wis Sandusisk River near Budryns, Ohio Sandusky River near Upper Sandusky, Ohio. Sandusky River near Freemont, Ohio.	Little Cuyahoga River at Akron, Ohio. Buffalo Creek at Gardenville, N. Y. Cayuga Creek near Lancaster, N. Y. Cazenovia Creek at Ebenezer, N. Y. Little Tonawanda Creek at Linden, N. Y.	Genesee River at Scio, N. Y. Genesee River at St. Helena, N. Y. Genesee River at Jones Bridge, near Mount	Canaseraga Creek near Dansville, N. Y	Bast Branch of Fish Creek at Taberg, N. Y. Black River near Boonville, N. Y. Black River at Waterform, N. Y. Moose River at McKeever, N. Y. Independence River at Sperryville, N. Y.	Oswegatchie River near Heuvelton, N. Y	Grass River at Pyrites, N. Y. Raquette River at Piercefield, N. Y. St. Regis River at Brasher Center, N. Y.
	No.		3-321	4-50. 4-119. 4-120.	4-134 4-139.5 4-140.3 4-140.7	4-1434-1444-146	4-148	4-180 4-197 4-201 4-204 4-209	4-229	4-230 4-233 4-236

53 51 64 140	22	8488	44 50 41 33 41	62 89 17
	-	5.36 1.38 11.12		0
		8.28	6.42	0
6.4.4.8.6 6.28.8.8	.57	. 87 4. 78 . 80 2. 86	. 02 0 7.10 0	0.0
1, 012 880 168 1, 630 524			620 550 760 770 725	
1,760 1,716 1,105 2,434 2,059	1,905	1, 191 914 1, 617 1, 329	1, 486 1, 656 1, 526 1, 005 1, 025	
3,355 3,830 5,112 5,344	5,344	4, 842 2, 727 4, 241 4, 241	2,923 4,135 3,315 1,420 1,362	1,362 1,441 840
27.8 25.5 47.1 21 54.9	30.5	49.8 35.8 42.7 71.0	15.7 19.8 32.7 14.6 20.7	88%
19.8 25.1 40.3 40.4	38.3	42.7 25.0 33.1 74.0	14.0 22.8 32.2 15.2 13.3	23.7 91.5 3.3
38. 63 27. 49 42. 4 36. 29 36. 96	47.2	24. 24. 44. 9. 9. 9	64.1 37.6 19.6 15.27 30.14	14. 56 3. 43 32. 9
132 218 122 413 351	433	271 177 248 178	318 455 197 127 145	136 63.3 115
712 559 283 1,059 1,230	1, 538	i. 061 828 814 729	1,022 1,120 578 1,039 1,012	1, 090 876 475
675 523 279 1, 051 1, 211	1, 548	1, 027 690 717 621	1, 037 1, 150 1, 160 980	1,006
766 610 285 1,071 1,256	1, 524	1, 107 1, 017 949 881	1, 196 1, 230 651 910 1, 044	1, 149
1, 493 1, 757 5, 948 1, 242 9, 913	3,615	5.516 2.566 5.259 25.960	1, 649 2, 566 627 552	1,616 34,720 5.2
1.38 1.42 1.85 1.85	1.94	2, 08 1, 70 1, 79 1, 56	1.1.2.2 1.1.3.40 1.2.840 1.2.840	1. 72 1. 85 2. 63
132 112 247 116 448	198	275 186. 5 307 628	76.1 139.0 140 77.1	119 699 3.01
Salmon River at Chasm Falls, N. Y. Chateaugay, River near Chateaugay, N. Y. Great Chazy River at Perry Mills, N. Y. West Brandr of Ausable Rivernear Newman, N. Y. Ausable River near Ausable Fiver near Service.	East Branch of Ausable River at Ausable	FOOKS N. Y.  BOUGHE River at Willsboro, N. Y.  Poultney River below Fair Haven, Vt.  Otter Creek at Middlebury, Vt.	Dog River at Northfield Falls, Vt. Mad River near Moretown, Vt. Ciyde River at Newport, Vt. Ciyde Crosse River near Leon, Wis. Coon Creek at Coon Valley, Wis.	Coon Creek near Stoddard, Wis. Kickapoo River at Stenhen, Wis. Raiston Creek at Iowa City, Iowa
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